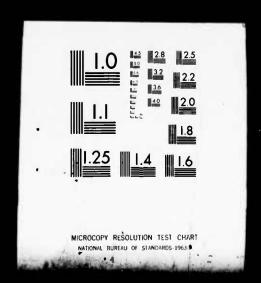


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### COLLAPSE OF TURBULENT WAKES IN STABLY STRATIFIED MEDIA

S. Hassid

Technical Memorandum File No. TM 77-237 10 August 1977 Contract No. N00017-73-C-1418

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Subject: Collapse of Turbulent Wakes in Stably Stratified Media

References: See page 22.

Abstract: A simple turbulent energy model and a calculation procedure

are proposed, which are able to reproduce the main features of the collapse of turbulent wakes behind towed or self propelled bodies in stratified environments. The predictions of the model are shown to be in good agreement with the detailed experimental data of Lin & Pao, as well as with the

flow visualization experiments of other workers.

Acknowledgment: This work was sponsored by the Naval Sea Systems Command,

Code NSEA-03133.



# Nomenclature

A,A',A",c <sub>1</sub> ,c <sub>2</sub> ,c <sub>1T</sub> ,	
c <sub>2T</sub> ,c <sub>T</sub> ,c <sub>ε1</sub> ,c <sub>ε2</sub> ,σ <sub>ε</sub>	constants in the model
c <sub>D</sub>	drag coefficient
D	body diameter
Fr .	Froude number (=U <sub>o</sub> /D N)
g	acceleration due to gravity
s <sub>i</sub>	acceleration due to gravity vector
н	wake height
k	turbulent energy per unit volume
N	Brunt - Väisälä frequency
P	mean pressure
Ps	pressure deviation from hydrostatic state
R	Richardson number $(R^{1/2} = r_0 \sqrt{\beta \alpha g/V_0})$
Rí	Richardson number (= $H^2N^2/k$ )
ro	initial radius of mixed region
Т	mean temperature
t <sub>col</sub>	time to start of collapse
v <sup>i</sup>	mean velocity vector
v <sub>d</sub>	velocity defect
$\mathbf{v}_{\mathbf{D}}$	velocity defect at the centerline
u <sub>o</sub>	body velocity
U,V,W	streamwise, horizontal cross-stream and vertical cross-stream mean velocity components
u <sup>1</sup>	fluctuating velocity vector
u,v,w	streamwise, horizontal cross-stream and vertical cross-stream fluctuating velocity component

u,v,w	root mean square fluctuating velocity components
W	wake width
x <sup>1</sup>	position vector
x,y,z	streamwise, horizontal cross-stream and vertical cross-stream coordinates, respectively
z	wake half-width
<sup>2</sup> 1/2	distance between centerline and point where $k = k(0)/4$
Greek Symbols	
α	temperature gradient of undisturbed fluid
β	volumetric expansion coefficient
ε	turbulent energy dissipation rate per unit volume
θ	temperature fluctuation
v <sub>v</sub> ,v <sub>w</sub>	eddy diffusivity coefficients for momentum in the horizontal and vertical direction, respectively
V <sub>Tv</sub> ,V <sub>Tw</sub>	eddy diffusivity coefficients for heat in the horizontal and vertical directions, respectively
ρ	density
Po	density at the centerline

### List of Figures

- Figure 1 Wake of a Self-Propelled Body in a Stably Stratified Fluid
  (a) Horizontal View, (b) Vertical View
- Figure 2 Mesh System, First Quadrant
- Figure 3 Nondimensional Width and Height of a Wake in Stratified and Nonstratified Media (a) Momentumless Wake, Data of Lin & Pao [3], (b) Drag Wake, Data of Pao & Lin [6]
- Figure 4 Height of Wakes in Stratified and Nonstratified Media
  (a) Momentumless Wake, Data of Lin & Pao [4], (b) Drag Wake,
  Data of Pao & Lin [7]
- Figure 5 Centerline Velocity Defect in Wakes in Stratified and Nonstratified Media (a) Momentumless Wakes, Data of Lin & Pao [4], (b) Drag Wakes, Data of Pao & Lin [7]
- Figure 6 Centerline Turbulent Energy in Wakes in Stratified and Nonstratified Media (a) Momentumless Wakes, Data of Lin & Pao [4], (b) Drag Wakes, Data of Pao & Lin [7]

### INTRODUCTION

A turbulent wake in a stably stratified fluid exhibits some peculiar features which are not encountered in wakes in nonstratified media.

Whereas axisymmetric wakes in homogeneous fluids increase indefinitely in width under the action of the diffusive turbulent forces, wakes in stably stratified environments are known to reach a maximum vertical extent and subsequently collapse to a minimum height. At the same time, under the combined action of turbulent diffusion and gravitational forces, they increase in the horizontal direction faster than they would in a neutral medium. Furthermore, since the decrease in the vertical extent is associated with a reduction in potential energy, gravitational internal waves are generated, which propagate through the surrounding fluid. (See Fig. 1.)

The first experimental investigation of this phenomena as consisted by Schooley & Stewart [1], who used a flow visualization as immique to observe the wake behind a self-propelled body. The same flow has since been documented in detail by Stockhausen, Clark & Kennedy [2], Lin & Pao [3], [4] and Stromm [5], whereas Pao & Lin [6], [7] focused on the wake behind a towed body in a stably stratified medium. Van der Watering, Tulin & Wu [8], Sumdaram et al. [9] and Meritt [10] investigated the collapse of mixed turbulent regions, generated by different means in quiescent, stably stratified media. These phenomena are closely related to the decay of momentumless wakes in such media, since turbulence in a momentumless wake is known to behave in a way similar to turbulence behind a point source. Wu [11] has observed the collapse of a tube of homogeneous, nonturbulent fluid in a stably stratified environment and the resulting wave pattern.

Theoretical investigations by Wessel [12] and Young & Hirt [13] focus on the collapse of a nonturbulent mixed region; this flow, however, differs from the collapse of the turbulent mixed region: whereas the latter reaches a final finite width, the former collapses indefinitely, until the fluid in the mixed region reaches the equilibrium level. The only theoretical treatment of the mixed region is by Vasiliev et al. [14]. The turbulence model used by these authors, however, does not predict any collapse, but only a slowing down of the rate of growth of the turbulent region. Furthermore, they neglect the nonlinear terms of the Navier Stokes equations. It is not clear that this approximation is justified because these are the very terms that cause the collapse. Finally, they make no comparison with experimental data and give no predictions for the streamwise velocity component in wake flows.

In this work it is proposed to use a simple turbulent flow model, coupled with a computational procedure, which will enable one to describe and document the phenomenon of collapse of turbulent wakes in stratified media.

### THE FUNDAMENTAL EQUATIONS

Consider the wake behind a body moving with uniform velocity in a stably stratified medium with uniform temperature (or density) gradient. (The application of the model to salinity-stratified media is straightforward). Assuming the Boussinesq approximation to be valid (i.e., that the specific weight of the fluid is variable but its density can be considered constant) one obtains the following form for the Reynolds stress and continuity equations

$$v_{j} \frac{\partial v_{i}}{\partial x_{j}} + \frac{1}{\rho} \frac{\partial P}{\partial x_{i}} = \beta g_{i} T - \frac{\partial}{\partial x_{j}} \overline{v_{i} v_{j}}$$
 (1)

$$\frac{\partial \mathbf{U}_{\mathbf{i}}}{\partial \mathbf{x}_{\mathbf{i}}} = 0 . (2)$$

In addition, the following simplifying assumptions will be made:

(a) The flow is elliptic in z and y and parabolic in x,

$$v_0 >> v_d , \qquad (3)$$

where  $\mathbf{U}_{o}$  is the velocity of the body and  $\mathbf{U}_{d}$  is the local velocity defect, and

(c) 
$$|\overline{vw}| \ll \overline{v^2}$$
 (4)

Thus the equations take the following form:

$$U_{o} \frac{\partial U_{d}}{\partial x} + V \frac{\partial U_{d}}{\partial y} + W \frac{\partial U_{d}}{\partial z} + \frac{\partial U^{2}}{\partial x} = -\frac{1}{\rho} \frac{\partial P_{s}}{\partial x} - \frac{\partial U_{v}}{\partial y} - \frac{\partial U_{w}}{\partial z} , \qquad (5)$$

$$U_{o} \frac{\partial V}{\partial x} + V \frac{\partial V}{\partial y} + W \frac{\partial V}{\partial z} + \frac{\partial V^{2}}{\partial y} = -\frac{1}{\rho} \frac{\partial P_{s}}{\partial y} , \qquad (6)$$

$$U_0 \frac{\partial W}{\partial x} + V \frac{\partial W}{\partial y} + W \frac{\partial W}{\partial z} + \frac{\partial W}{\partial z} = -\frac{1}{0} \frac{\partial P}{\partial z} - \beta g(T - \alpha z) , \qquad (7)$$

and

$$\frac{\partial V}{\partial y} + \frac{\partial W}{\partial z} = 0 . (8)$$

Here  $P_s$  is the deviation of the pressure from the unperturbed hydrostatic pressure, whereas  $\alpha z$  is the unperturbed temperature distribution, which is assumed linear.

By differentiating Eq. (6) with respect to y and Eq. (7) with respect to z and adding, a Poisson equation for the pressure can be obtained:

$$\frac{\partial^2}{\partial y^2} \left( \frac{P_s}{\rho} + \overline{v^2} \right) + \frac{\partial^2}{\partial z^2} \left( \frac{P_s}{\rho} + \overline{w^2} \right) = -\beta g \frac{\partial}{\partial z} \left( T - \alpha z \right) . \tag{9}$$

In addition, the following equation governs the transport of temperature:

$$U_{o} \frac{\partial T}{\partial x} + V \frac{\partial T}{\partial y} + W \frac{\partial T}{\partial z} = -\frac{\partial}{\partial y} \overline{v\theta} - \frac{\partial}{\partial z} \overline{w\theta} . \qquad (10)$$

In order to solve the above system of equations, closure approximations are needed for the Reynolds stresses  $\overline{uv}$  and  $\overline{uw}$ , as well as for the turbulent heat fluxes  $\overline{v\theta}$  and  $\overline{w\theta}$ . These will be discussed in the following section.

### THE TURBULENT FLOW MODEL

Recently, the Reynolds stress model has become increasingly popular in computational investigations of turbulent flow. For flows in wakes behind both towed and self-propelled bodies, however, the simpler  $k-\epsilon$  model gives results which are in better agreement with experiment. (See Hassid [15]). Needless to say, of course, the  $k-\epsilon$  model is also preferable because of its lower requirements in computer storage and time.

The closure used here is essentially based on the model of McGuirk and Rodi [16]. Transport equations are used for both the turbulent energy k and the dissipation rate  $\epsilon$  and gradient diffusion forms for the Reynolds stresses. However, the expressions for the latter quantities are derived from invariant second order models by neglecting the advection and turbulent transport terms.

The expressions for  $u_1u_j$  and  $u_1\theta$  are

$$\frac{\overline{\mathbf{u_i u_j}}}{k} = \frac{(1-c_2)}{c_1} \frac{\overline{\mathbf{I_{ij}}}}{\varepsilon} - \frac{2}{3} \frac{(1-c_1-c_2)}{c_1} \delta_{ij}$$
(11)

where

$$\Pi_{\mathbf{i}\mathbf{j}} = -\left(\overline{u_{\mathbf{i}}u_{\mathbf{k}}} \frac{\partial U_{\mathbf{j}}}{\partial x_{\mathbf{k}}} + \overline{u_{\mathbf{j}}u_{\mathbf{k}}} \frac{\partial U_{\mathbf{i}}}{\partial x_{\mathbf{k}}}\right) - \left(g_{\mathbf{i}}\overline{u_{\mathbf{j}}\theta} + g_{\mathbf{j}}\overline{u_{\mathbf{i}}\theta}\right) \tag{12}$$

$$\overline{\mathbf{u}_{\underline{i}}\theta} = -\frac{\mathbf{k}}{\mathbf{c}_{\underline{1}\underline{T}}\varepsilon} \left[ \overline{\mathbf{u}_{\underline{i}}\mathbf{u}_{\underline{k}}} \frac{\partial \underline{\mathbf{T}}}{\partial \mathbf{x}_{\underline{k}}} + (1 - \mathbf{c}_{\underline{2}\underline{T}}) (\overline{\mathbf{u}_{\underline{k}}\theta} \frac{\partial \underline{\mathbf{U}}_{\underline{i}}}{\partial \mathbf{x}_{\underline{k}}} + \beta \ \mathbf{g}_{\underline{i}} \overline{\theta^{2}}) \right] . \tag{13}$$

 $\theta^2$ , too, is calculated assuming equilibrium between production and dissipataion of this quantity. It is also assumed that the rate of destruction of the temperature variance  $\theta^2$  is proportional to the mechanical dissipation

$$\overline{\theta^2} = -\frac{2}{c_T} \frac{k}{\varepsilon} \overline{u_j \theta} \frac{\partial T}{\partial x_j} . \qquad (14)$$

McGuirk and Rodi used their model to describe jet flows in stratified media. Wake flows differ from jet flows in that the latter are parabolic in y. Therefore, in our case, some further approximations are needed.

As already stated, ww will be neglected in comparison to the normal Reynolds stress components. Thus, the equations for the turbulent heat fluxes are

$$-\overline{u\theta} = \frac{k}{c_{1T}} \left[ \overline{uw} \frac{\partial T}{\partial z} + \overline{uv} \frac{\partial T}{\partial y} + (1 - c_{2T}) \left( \overline{w\theta} \frac{\partial U}{\partial z} + \overline{v\theta} \frac{\partial U}{\partial y} \right) \right]$$
(15)

$$-\overline{v\theta} = \frac{k}{c_{1T}\varepsilon} \overline{v\theta} \frac{\partial T}{\partial y}$$
 (16)

$$-\overline{w\theta} = \frac{k}{c_{1T}\varepsilon} \left[ \overline{w^2} \frac{\partial T}{\partial z} - (1-c_{2T}) \beta g \overline{\theta^2} \right] . \qquad (17)$$

These equations will be further simplified by assuming that  $\partial T/\partial y$  is much smaller than  $\partial T/\partial z$ . Although this assumption is not strictly true, the error caused will be small, both for weak stratification, because the dependence on  $\partial T/\partial y$  is small in that case, and for strong stratification, because then  $\partial T/\partial z$  is indeed much larger than  $\partial T/\partial y$ . Thus Eq. (14) becomes

$$\overline{\theta^2} = -\frac{2}{c_T} \frac{k}{\varepsilon} \overline{w\theta} \frac{\partial T}{\partial z} . \qquad (18)$$

Substituting Eq. (18) into Eq. (17) one finds

$$\overline{w\theta} = \frac{\frac{k \overline{w^2}}{c_{1T} \varepsilon} \frac{\partial T}{\partial z}}{1 + \frac{(1 - c_{2T})}{c_{T}} \frac{k^2}{\varepsilon^2} \beta g \frac{\partial T}{\partial z}} .$$
 (19)

Now, from Eq. (11), one obtains the following expressions for  $v^2$  and  $w^2$ 

$$\overline{v^2} = -\frac{2}{3} \frac{(1-c_1-c_2)}{c_1} k \qquad , \tag{20}$$

$$\frac{1}{w^2} = -\frac{w}{3} \frac{(1-c_1-c_2)}{c_1} k - 2 \beta g \frac{k}{\varepsilon} \overline{w\theta} . \qquad (21)$$

Substituting into Equations (16) and (19) one obtains

$$-\overline{v\theta} = v_{Tv} \frac{\partial T}{\partial y} = c \frac{k^2}{\varepsilon} \frac{\partial T}{\partial y} , \qquad (22)$$

$$-\overline{w\theta} = v_{Tw} \frac{\partial T}{\partial z} = \frac{c_{\theta} \frac{k^2}{\varepsilon} \frac{\partial T}{\partial z}}{1 + A \frac{k^2}{\varepsilon^2} \beta g \frac{\partial T}{\partial z}},$$
 (23)

where

$$c_{\theta} = -\frac{2}{3} \frac{(1-c_1-c_2)}{c_1 c_T}$$
, (24)

and

$$A = \frac{2}{c_{1T}} \left[ \frac{1 - c_{2T}}{c_{T}} + \frac{1 - c_{2}}{c_{1}} \right] \qquad (25)$$

Turning to the equations for the Reynolds stress one has

$$\overline{uv} = -\frac{1-c_2}{c_1} \frac{k v^2}{\varepsilon} \frac{\partial U}{\partial y} , \qquad (26)$$

$$\overline{uw} = -\frac{(1-c_2)}{c_1} \left[ \frac{k \overline{w^2}}{\varepsilon} \frac{\partial U}{\partial z} - 2 \beta g \overline{u\theta} \right] . \qquad (27)$$

Again, using Equations (20) and (26) one obtains

$$-\overline{uv} = v_v \frac{\partial U}{\partial y} = c_M \frac{k^2}{\varepsilon} \frac{\partial U}{\partial y} , \qquad (28)$$

where

$$c_{M} = -\frac{2}{3} \frac{(1-c_{2})}{c_{1}} \frac{(1-c_{1}-c_{2})}{c_{1}} . (29)$$

Neglecting, as already mentioned,  $\partial T/\partial y$  in the equation for  $u\bar{\theta}$  and substituting in Eq. (27), one ends with

$$-\overline{uw} = v_W \frac{\partial U}{\partial z} = c_M \frac{k^2}{\varepsilon} \frac{(1+A'F)}{(1+A'F)(1+AF)} \frac{\partial U}{\partial z} , \qquad (30)$$

where

$$F = \beta g \frac{k^2}{\epsilon^2} \frac{\partial T}{\partial z} , \qquad (31)$$

and

$$A' = A - \left(2 \frac{1-c_2}{c_1} + \frac{c_{2T}}{c_{1T}}\right) \frac{1}{c_{1T}}$$
, (32a)

$$A'' = \frac{1-c_2}{c_{1T} c_1} . (32b)$$

As already mentioned, k and  $\epsilon$  are calculated from the relevant transport equations:

$$U_{o} \frac{\partial k}{\partial x} + V \frac{\partial k}{\partial y} + W \frac{\partial k}{\partial z} = \frac{\partial}{\partial y} \left[ v_{v} \frac{\partial k}{\partial y} \right] + \frac{\partial}{\partial z} \left[ v_{w} \frac{\partial k}{\partial z} \right] + v_{v} \left( \frac{\partial U}{\partial y} \right)^{2}$$

$$+ v_{w} \left( \frac{\partial U}{\partial z} \right)^{2} + \beta g \overline{w\theta} - \varepsilon , \qquad (33)$$

$$U_{o} \frac{\partial \varepsilon}{\partial x} + V \frac{\partial \varepsilon}{\partial y} + W \frac{\partial \varepsilon}{\partial z} = \frac{\partial}{\partial y} \left[ \frac{v_{v}}{\sigma_{\varepsilon}} \frac{\partial \varepsilon}{\partial y} \right] + \frac{\partial}{\partial z} \left[ \frac{v_{w}}{\sigma_{\varepsilon}} \frac{\partial \varepsilon}{\partial z} \right] - c_{\varepsilon 2} \frac{\varepsilon^{2}}{k}$$

$$+ c_{\varepsilon 1} \frac{\varepsilon}{k} \left[ v_{v} \left( \frac{\partial U}{\partial y} \right)^{2} + v_{w} \left( \frac{\partial U}{\partial z} \right)^{2} + \beta g \overline{w\theta} \right] . \quad (34)$$

Here the model differs from the one of McGuirk and Rodi because these authors claim that gravitational production does not appear in the equation for  $\epsilon$ . It was found that such an assumption leads to unrealistically low values for k. It should be mentioned that there is wide an greement among different research groups on the question of the influence of buoyant forces on the  $\epsilon$  equation.

The constants used are

c <sub>1</sub>	c <sub>2</sub>	c <sub>E</sub> 1	c <sub>E2</sub>	σε	c <sub>1T</sub>	<sup>C</sup> 2T	c <sub>T</sub>
2.2	.55	1.44	1.92	1.3	3.2	.5	1.25

Of all the approximations used the neglect of  $\partial T/\partial y$ , which was necessary in order to obtain the gradient forms for the Reynolds stress equations, is the one likely to cause the greatest error. Notice however,

that for the non-buoyant case, the expressions derived above are exact.

### METHOD OF SOLUTION

The system of equations to be solved consists of the x and y momentum equations

$$U_{o} \frac{\partial U_{d}}{\partial x} + V \frac{\partial U_{d}}{\partial y} + W \frac{\partial U_{d}}{\partial z} + \frac{\partial}{\partial x} (\frac{P_{s}}{\rho} + V^{2}) + \frac{\partial}{\partial x} (\overline{u^{2}} - \overline{v^{2}}) =$$

$$\frac{\partial}{\partial y} \left[ (v_{v} \frac{\partial U_{d}}{\partial y}) + \frac{\partial}{\partial z} \left[ (v_{w} \frac{\partial U_{d}}{\partial z}) \right]$$
(35)

$$U_{o} \frac{\partial V}{\partial x} + V \frac{\partial V}{\partial y} + W \frac{\partial V}{\partial z} + \frac{\partial}{\partial y} \left( \frac{P_{s}}{\rho} + \overline{V^{2}} \right) = 0 , \qquad (36)$$

the transport equation for temperature

$$U_{o} \frac{\partial T}{\partial x} + V \frac{\partial T}{\partial y} + W \frac{\partial T}{\partial z} = \frac{\partial}{\partial y} \left( v_{Tv} \frac{\partial T}{\partial y} \right) + \frac{\partial}{\partial z} \left( v_{Tw} \frac{\partial T}{\partial z} \right) , \qquad (37)$$

the Poisson equation for pressure

$$\frac{\partial^{2}}{\partial y^{2}} \left( \frac{P_{s}}{\rho} + \overline{v^{2}} \right) + \frac{\partial^{2}}{\partial z^{2}} \left( \frac{P_{s}}{\rho} + \overline{w^{2}} \right) + \frac{\partial^{2}}{\partial z^{2}} \left( \overline{w^{2}} - \overline{v^{2}} \right) = \beta g \frac{\partial}{\partial z} \left( T - \alpha z \right) , \qquad (38)$$

the continuity equation (Eq. 8), the transport equations for the turbulent energy k (Eq. 33), and the dissipation rate  $\varepsilon$  (Eq. 34). In addition, the following expressions for  $u^2-v^2$  and  $w^2-v^2$  can be derived from Eq. (11)

$$\frac{\overline{u^2} - \overline{v^2}}{\overline{u^2} - \overline{v^2}} = -\frac{(1 - c_2)}{c_1} \left( \overline{uv} \frac{\partial U}{\partial y} + \overline{uw} \frac{\partial U}{\partial z} \right)$$
(39)

and

$$\frac{\overline{v^2} - \overline{v^2}}{v^2 - v^2} = \frac{(1 - c_2)}{c_1} \beta g \overline{w\theta}$$
 (40)

This system was converted into a set of difference equations. Forward differencing was used in the x direction, centered differencing for the diffusive terms and upwind differencing for the lateral transport terms.

The mesh system used (Fig. 2) consists of an inner mesh AEFG in which the turbulent processes occur and an outer mesh, in which there is no turbulence. The outer mesh was three times as large as the inner mesh. The reason is that both experimental evidence and computer calculations suggest that the waves caused by the collapse of a nonturbulent mixed region advance twice the original width of that region after two Brunt-Väisälä periods. Since there are very few experimental data for times larger than two Brunt-Väisälä periods and in addition, the wave phenomena are much weaker when the mixed region is turbulent, this mesh was estimated to be adequate.

The mesh used was allowed to expand, so that the original growth due to turbulent diffusion can be simulated.

The following boundary conditions were used:

On AB, 
$$v = \frac{\partial W}{\partial y} = \frac{\partial}{\partial y} \left( \frac{P_s}{\rho} + \overline{v^2} \right) = \frac{\partial U_d}{\partial y} = 0 .$$
 (41)

On AD, 
$$\frac{\partial V}{\partial z} = W = \frac{\partial}{\partial z} \left( \frac{P_s}{\rho} + \overline{W^2} \right) = \frac{\partial U_d}{\partial z} = 0 \qquad (42)$$

On BC, 
$$V = \frac{\partial W}{\partial z} = P_s = U_d = 0 . \tag{43}$$

On CD, 
$$\frac{\partial V}{\partial y} = W = P_s = U_d = 0 . \qquad (44)$$

For k and  $\epsilon$ , the boundary conditions were

$$k = \varepsilon = 0 . (45)$$

$$\frac{\partial \mathbf{k}}{\partial \mathbf{y}} = \frac{\partial \varepsilon}{\partial \mathbf{y}} = 0 . (46)$$

$$\frac{\partial \mathbf{k}}{\partial z} = \frac{\partial \varepsilon}{\partial z} = 0 . (47)$$

Admittedly, some of these conditions are artificial and may result in errors as the wave fronts approach the mesh boundaries. It is estimated, however, that these errors will not be excessive for times less than two Brunt-Väisälä periods. As a check, the differential equations were solved using different boundary conditions and this did not significantly affect the results in the turbulent region.

### RESULTS AND COMPARISON WITH EXPERIMENT

The main scaling parameter used by most investigators (Refs. 3, 6, 8, 9 and 10) for wake flows in stratified media is the Froude number, defined as

$$Fr = \frac{U_O}{N D} \tag{48}$$

where N is the Brunt-Väisälä frequency

$$N = \frac{1}{2\pi} \left( \frac{g}{\rho} \frac{\partial \rho}{\partial y} \right)^{1/2} \qquad . \tag{49}$$

However, this scaling factor is not sufficient; what determines the extent that a momentumless wake will grow before collapsing is rather the Richardson number, which expresses the ratio between the potential energy and the turbulent kinetic energy at the origin

$$Ri = \frac{H^2 N^2}{k} \qquad . \tag{50}$$

The decay of momentumless wakes, even in the self-similar state, depends on such factors as the body shape (whether the body is blunt or slender), the propelling mechanism, etc. For non-buoyant flows behind slender bodies, the wake width is given by

$$\frac{H}{D} = .8 \left(\frac{x}{D}\right)^{.22} \qquad . \tag{51}$$

(See Lin & Pao [3]). In Fig. 3a the influence of stratification on the wake behind such a body is shown. It is seen that the wake height and width become functions of  $Nx/U_0$ . Collapse starts when  $Nx/U_0 = 1/3$ . The wake height decreases by 10% and then increases slightly. The wake width increases throughout the flow.

In the same figure, the correlations of Lin & Pao [3] are shown. The agreement between the predicted and the measured wake width is very good; however, our predictions overestimate the height of the wake. It must be noted, however, that Lin & Pao concede that there is considerable difficulty in estimating the wake width because of the irregular fluctuating nature of the outer wake boundary.

In Fig. 3b the width and the height of the wake behind a towed body in stratified flow are shown. Here the computed values of  $2z_{1/2}$  are compared with the values of H and W measured by Pao & Lin [5] who claim that for drag wakes H =  $2z_{1/2}$ . (In our computations, H =  $2.4z_{1/2}$ .) It is seen that in wakes the first maximum is reached at a higher value of Nt than in momentumless wakes. Though the computations fail to reproduce the 10%

reduction in width recorded in the experiments, the agreement can be considered satisfactory, in view of the influence of stratification on the initial wake widths, which cannot be accounted for by the calculation procedure and in view of the experimental errors of flow visualization.

In Figs. 4a and 4b, the predicted wake width is compared with the results of Lin & Pao [4] for momentumless wakes and Pao & Lin [7] for drag wakes. The system of maxima and minima observed in these experiments is reproduced; there is however, a certain difference in phase probably due to the use of V=O and W=O as initial conditions.

In Fig. 5a and 5b it is seen that stratification results in slower decay of the velocity defect, both for drag wakes the momentumless wakes. Agreement with the experimental data is very good.

In Figs. 6a and 6b the decay of turbulent energy with and without stratification is shown. It is seen that in spite of the additional destruction of turbulent energy due to the stable stratification the difference in the turbulence levels is very small, especially in the momentumless wake. Agreement between theory and experiment is again very good in spite of the fact that the predicted  $\sqrt{k}$  is compared with the measured  $\overline{u}$ .

Finally in Table I the predicted values of the maximum and the minimum heights as well as the collapse times of mixed regions and momentumless wakes are compared with their values as measured by several experimental workers.

In this table,  $t_{col}$  is the time (for mixed regions; distance divided by mean body velocity in momentumless wakes) at which the wake reaches the first maximum height,  $Z_{max}$  and  $Z_{min}$  are the maximum and minimum wake heights, respectively, whereas the Richardson number R is defined as

$$R^{1/2} = \frac{r_o \sqrt{\beta \alpha g}}{V_o} \qquad (52)$$

Table I

Collapse Time, Maximum Height and Minimum Height of Turbulent Mixed Regions in Stably Stratified Media

Referenc	e		t <sub>co1</sub>	βag	Z max	/r <sub>o</sub>	Z <sub>min</sub>	/r <sub>o</sub>
			Meas.	Pred.	Meas.	Pred.	Meas.	Pred.
Schooley & Stewart	Fr =	60	2.12	1.8	2.16	2.00	1.36	1.8
Stockhausen, Clark								
& Kennedy	Fr =	100	1.63	1.2	3.00	2.08	2.10	1.8
Van der Watering,	1/0							
Tulin & Wu	$R^{1/2}$	= 1.34	1.33	.92	1.51	1.26	1.11	.96
		1.56	1.48	.98	1.25	1.23	1.04	.91
		1.56	1.21	.98	1.31	1.26	1.06	.93
		2.19	1.35	.90	1.11	1.17	. 90	.80
		2.19	1.35	.91	1.19	1.17	90	.82
		3.09	1.40	.78	1.06	1.09		.68
		4.23	1.23	.51	.95	1.03	.77	.57
		4.27	1.05	.69	1.00	1.05	.83	. 62
		1.56	1.12	.93	1.24	1.20	.94	. 86
		1.57	1.12	.94	1.25	1.21	.89	. 87
		2.28	.89	.74	1.20	1.13	1.00	.72
		3.21		.93	1.13	1.13	.94	.63
		2.06	1.05	.65	1.20	1.06	1.03	.63
		1.89	.73	.70	1.04	1.07	.88	.65
		3.69	1.07	.50	1.04	1.09		.58
		5.21		.58	.88	1.04	.77	. 59
		5.29	1.18	.45	. 95	1.02	.88	.55
Sundaram, Straton		3.20	3.75	.72	1.54	1.12	1.07	.71
& Rehm		2.80	2.40	.77	1.31	1.13	1.00	.74

where  $V_{o}$  is the initial rate of spreading and  $r_{o}$  is the initial radius of the mixed region. For the experiments of Schooley & Stewart and Stockhausen, Clark & Kennedy,  $r_{o}$  is the body diameter.

There are of course several sources of error in this table. Since no initial turbulent energy data were available, this quantity was deduced from the initial rate of spread and the equation

$$\frac{1}{\sqrt{k}}\frac{d}{d}\frac{Z}{t} = .26 \tag{53}$$

which is applicable to self-similar flow of momentumless wakes in nonstratified media. The initial turbulent energy and dissipation profiles were also assumed self-similar. Given, however, the very high Richardson numbers, it is very unlikely that the turbulent energy distribution may reach self similarity. Besides, the predicted value of t<sub>col</sub> is the value of t when the wake height actually starts decreasing; typically, however, it has decreased by less than 5% at t=2 t<sub>col</sub> and it is only then that it can be observed in a flow visualization experiment. Finally, for the experiments of Schooley & Stewart and Stockhausen, Clark & Kennedy, the production of turbulence by mean velocity gradients could not be taken into account.

In spite of all these considerations, it is thought that in view of the crudeness of the observation methods, the numerous assumptions needed for the calculation and the disagreement between the data themselves, even the order of magnitude agreement presented in Table I is satisfactory.

### DISCUSSION AND CONCLUSIONS

The development of momentumless wakes, drag wakes and turbulent regions in stably stratified media was investigated using a simple turbulent energy - dissipation model. This is the simplest model that can be used for such flows; evidence from calculations in nonstratified media suggests

that for wake flows, this model is also more accurate than the more sophisticated Reynolds stress models.

The effect of stratification is expressed through a damping factor in the eddy diffusivity coefficients for vertical exchange of momentum and heat. The eddy diffusivities in the horizontal direction remain unaffected.

The stabilizing temperature gradient acts in two ways: (a) by reducing the turbulent fluctuation level in the vertical direction, thus impeding the ability of turbulence to diffuse in this direction; (b) by generating a mean flow pattern in which the vertical velocity is directed towards the horizontal symmetry plane and the horizontal velocity is directed away from the vertical symmetry plane. This flow pattern is due to the difference in weight between the fluid in the turbulent region and the surroundings. Thus, expansion of the mixed regions is impeded in the vertical direction and enhanced in the horizontal direction. The wake eventually reaches a maximum height in the vertical direction and shrinks. A system of maxima and minima in the height follows, which are separated by roughly one Brunt-Väisälä period.

The stratification is shown to slow the decay in the velocity defect, presumably as a result of the decrease in the rate of vertical exchange of momentum. It has very little effect on the total turbulent energy, because the destruction of turbulent energy by gravitational forces is balanced by the increase in the production (since the rate of decay of the mean velocity is smaller) and the decrease in the diffusion. These features of the flow, which are well documented in the experiments of Lin & Pao, are well reproduced by the model.

The model has, however, several limitations. It is not known whether it can account for the laminarization, that Lin & Pao claim occurs after four Brunt-Väisälä periods. The influence of gravitational negative production of turbulent energy on the dissipation equation is not certain and the particular assumption used in this model is only a tentative one, to be modified when a more realistic model is available. The calculation procedure used is not reliable for times larger than two Brunt-Väisälä periods, because at such times the artificially imposed boundary conditions may result in reflection of the waves generated by the collapse. In order to overcome this drawback, one may introduce a region of high viscosity on the outer part of the flow, so that outcoming waves are damped before they can be reflected. If this is done one may have a more reliable view of the wave pattern caused by the collapse but the computer requirements in storage and time would be significantly increased. Finally one should mention the artificial viscosity errors caused by the upwind differencing technique used to approximate the lateral transport terms which may partially account for the fact that the predicted collapse rate is lower (i.e., the final thickness is higher) than in many (but not all) experiments.

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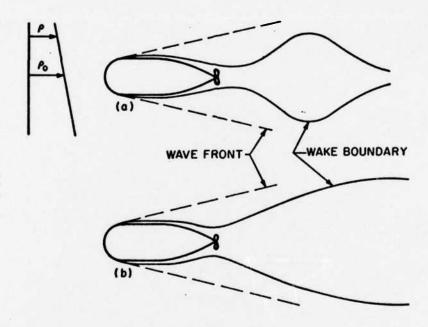


Figure 1 - Wake of a Self-Propelled Body in a Stably Stratified Fluid (a) Horizontal View, (b) Vertical View

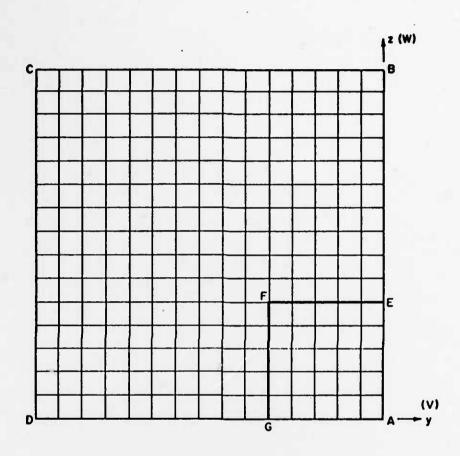
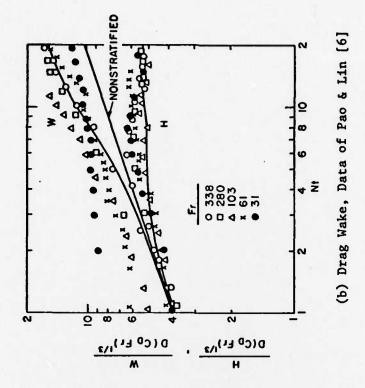


Figure 2 - Mesh System, First Quadrant



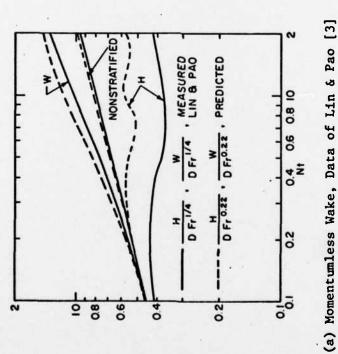
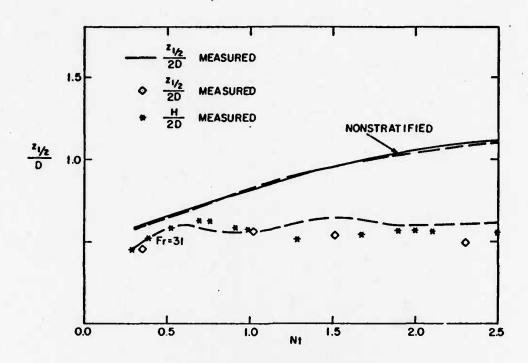
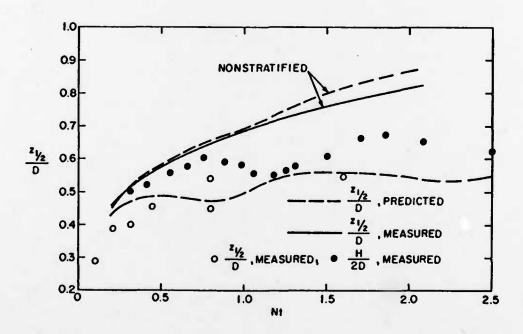


Figure 3 - Nondimensional Width and Height of a Wake in Stratified and Nonstratified Media

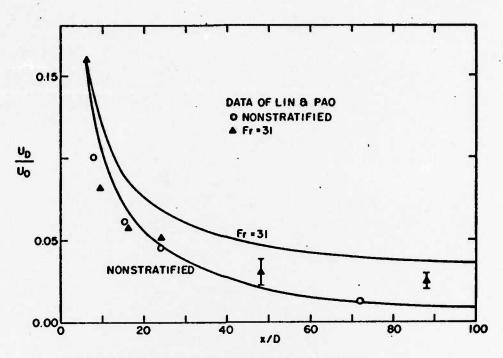


(a) Momentumless Wake, Data of Lin & Pao [4]

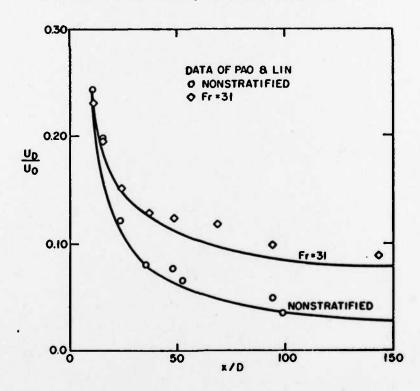


(b) Drag Wake, Data of Pao & Lin [7]

Figure 4 - Height of Wakes in Stratified and Nonstratified Media



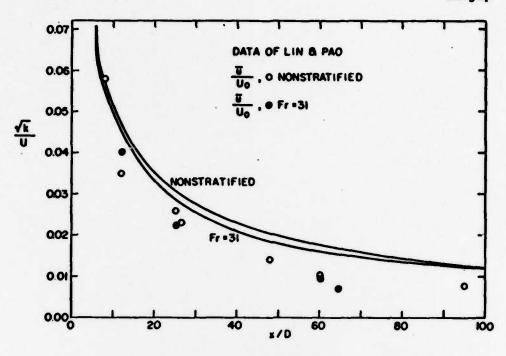
(a) Momentumless Wakes, Data of Lin & Pao



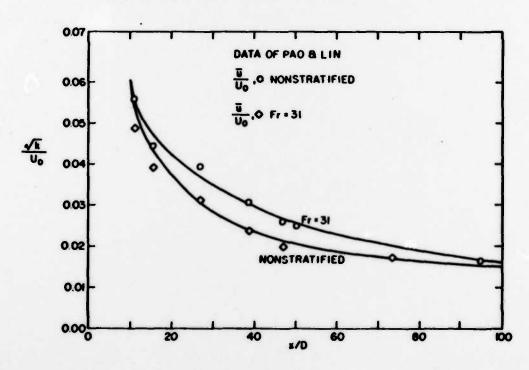
(b) Drag Wakes, Data of Pao & Lin [7]

)

Figure 5 - Centerline Velocity Defect in Wakes in Stratified and Nonstratified Media



(a) Momentumless Wakes, Data of Lin & Pao [4]



(b) Drag Wakes, Data of Pao & Lin [7]

Figure 6 - Centerline Turbulent Energy in Wakes in Stratified and Nonstratified Media

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